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Superstructures, (Si,Al) and H₂O ordering in armenite BaCa₂Al₆Si₉O₃₀ · 2 H₂O*

by Thomas Armbruster¹

Abstract

Armenite is a member of the group of double-ring silicates. Partial (Si,.Al) ordering leads to orthorhombic symmetry which can be resolved with $FeK\alpha_1$ Guinier powder photographs. Using single-crystal diffraction experiments superstructure reflections caused by H_2O ordering indicate a primitive lattice with a = 13.874(2), b = 18.661(2), and c = 10.697(1) Å. The crystal structure of an optically homogeneous armenite from Rémigny (Canada) was determined and refined from X-ray single-crystal data in the space group Pnna.

Keywords: Armenite, ordering, superstructure, structure refinement, X-ray methods.

Introduction

Armenite (BaCa₂Al₆Si₉O₃₀ · 2 H₂O) together with milarite and osumilite belong to the group of double-ring silicates. This rare mineral was detected and described in several locations including: the Armen Mine, Norway (Neumann, 1941). Rémigny, Canada (Pouliot et al., 1984), Broken Hill, Australia (Mason, 1987), Zufuru Mine in Sardinia, Italy (Balassone et al., 1989) and the celsian-bearing leucocratic gneisses (Frank, 1979) of the Simplon area in Switzerland (Senn, 1990).

Armenite when viewed under a polarizing microscope shows pseudo-hexagonal morphology formed by penetration twins with domains showing an 120° angle (Neumann, 1941; Pouliot et al., 1984), Neumann (1941) recognized armenite as a biaxial, $2 V_x = 60^\circ$ ("orthorhombic?") mineral. Bakakin et al. (1975) were aware of this twinning problem, but refined the structure of an armenite twin-aggregate in the space-group P6/mcc as is characteristic of other double-ring silicates. Since this study, armenite has erroneously been considered hexagonal.

A series of studies on double-ring silicates was conducted to study the relation between optical anomalies and symmetry reduction. Thin sections of milarite (KNa_{1-x} Ca₂Be_{3-x}Al_xSi₁₂O₃₀ · [H₂O])

with the space group P6/mcc revealed optical sectors in (001) and were related to partial ordering of H₂O (ARMBRUSTER et al., 1989). This study reports experiments performed on armenite from Rémigny Canada, which was kindly provided by the Museum of Natural History Bern, sample NHMB-B4468. A more detailed description of these structural investigations will be published elsewhere.

Mineralogical studies

X-ray powder diffraction of the armenite sample from Rémigny, Canada, (recorded on a Guinier camera with FeK α_1 -radiation) displays characteristic splitting of various reflections (Tab. 1). The reflections can not be indexed with a hexagonal lattice but are in agreement with a face-centered orthorhombic lattice (k + l = 2 n). Pouliot et al. (1984) previously reported a powder diffraction pattern determined using a diffractometer with CuK α -radiation but could not resolve any splitting.

A thin section of an armenite twin was prepared perpendicular to the pseudo-hexagonal caxis and an optically homogeneous domain was separated with the micro-drilling device of Medenbach (1986). This domain was studied with

^{*}Extended abstract of a contribution presented at the annual meeting of the SSMP, Geneva, October 4 and 5, 1990.

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Tab. 1 Indexed powder pattern of natural orthorhombic (Pnna) armenite from Rémigny Canada: a=13.874(2), b=18.661(2), c=10.697(1) Å and of dehydrated armenite (Amaa): a=14.024(3), b=18.383(3), c=10.606(2) Å. Intensities were calculated (Yvon et al., 1977) for a Guinier camera (FeK α_1 -radiation) using the positional parameters obtained from the structure refinement.

n	atur	al	armenite			dehydrated	armenite
h	k	1	d _{calc}	d _{obs} 1	calc	^d calc	d _{obs}
0	2	0	9.330	9.291	23	9.191	9.195
0		1	9.280	11	34	9.188	н
2	0	0	6.937	6.927 1	00	7.012	7.008
0		1	5.377	5.388	7	5.306	5.307
0	0	2	5.349	11	4	5.304	11
2	3,	1	4.2498	4.255	50	4.2312	4.2307
2	0	2	4.2357	4.247	24	4.2301	μ
2	4	0	3.8711	3.8712	31	3.8437	3.8404
2	2	2	3.8569	3.8633	66	3.8427	0.4706
0	1	3	3.5023	0 4604	2	3.4723	3.4726
4	0	0	3.4685	3.4684	8	3.5060	3.5063
1		1	3.4153	3.4182	29 33	3.3720 3.3714	3.3719
1		2	3.4079 3.3958	3.4091 3.3961	28	3.3706	11
0	6	0	3.1100	3.1098	14	3.0638	3.0609
0		3	3.0934	3.0942	30	3.0627	"
4		1	2.9147	2.9148		2.9251	2.9246
4	o	2	2.9101	2.5110	31	2.9248	11
3	4	2	2.7988	2.7986	6	2.7880	
4	4	0	2.7834	2.7841	28	2.7875	2.7885
4	2	2	2.7781	2.7782	49	2.7871	11
0	6	2	2.6885	2.6894	28	2.6530	2.6557
0	0	4	2.6742	2.6749	15	2.6520	2.6481
0	7	1	2.5866	2.5873	7	2.5492	2.5495
0	2	4	2.5707	2.5717	6	2.5481	
1	7	1	2.5418	2.5433	6	2.5081	2 5001
1	5	3	2.5347	2.5336	5	2.5075	2.5081
1 2	2	2	2.5280	2.5272	5 18	2.5070 2.4814	2.4825
2	0	4	2.5072 2.4955	2.5066	9	2.4805	2.4779
4	5	1	2.4722	2.4727	5	2.4677	2.4681
4	4	2	2.4694	2.4687	4	2.4675	0
O	4	4	2.3204	2.3190	1	2.2970	
4	6	0	2.3155	2.3147	3	2.3070	2.3075
4	3	3	2.3089	2.3081	4	2.3065	11
3	5	3	2.2518		2	2.2376	2.2371
3	2	4	2.2469		.2	2.2373	"
6	2	0	2.2444	121 1250200	1	2.2652	2.2640
2	8	0	2.2109	2.2104	4	2.1836	2.1826
2	4	4	2.2006	2.2006	8	2.1829	
6	3	1	2.1242		9	2.1390	2.1391
6	0	2	2.1225	2 1172	4	2.1389 2.1151	2.1152
6	0 4	0	2.1181 2.0718	2.1172	2	2.0834	2.0818
6	2	2	2.0696		4	2.0832	"
2	8	2	2.0435	2.0430	4	2.0192	2.0200
0	9	1	2.0355	2.0348		2.0057	2.0052
0	6	4	2.0277	2.0266	7	2.0052	
2	9	1	1.9531	1.9527	10	1.9283	1.9282
2	6	4	1.9465	1.9474	9	1.9279	"
5	7	1	1.8920	1.8929	8	1.8864	1.8860
5	5	3	1.8889	1.8882	8	1.8862	"
0	10	0	1.8660		2	1.8384	1.8366
0	5	5	1.8561		7	1.8376	
6	6	0	1.8556		4	1.8583	1.8581
6	3	3	1.8522	1.8517	8	1.8581	
6	6	2	1.7533	1.7537	18 7	1.7536 1.7535	1.7539
6	0	4	1.7493	1.7496		1.7530	
8	0	U	1.7343	1.1339	12	1.7550	

a spindle-stage equipped polarizing microscope and extinction data yielded an optic axial angle 2 $V_x = 59.9(6)$. This result is in excellent agreement with the U-stage measurements of Neumann (1941) on armenite from Norway. The same isolated single-crystal from Rémigny was transferred to a precession camera (MoK α -radiation)

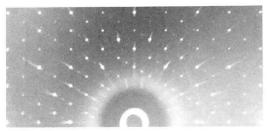


Fig. 1 hk2-layer (pseudo-hexagonal setting) of an optically homogeneous armenite from Rémigny, Canada (recorded with a precession camera, MoKα-radiation). In addition to the pseudo-hexagonal arrangement of reflections, slightly streaked superstructure reflections which characterize a primitive orthorhombic lattice can be observed. Systematic extinctions lead to space group Pnna

where the hk2 layers (pseudo-hexagonal setting) displayed superstructure reflections (Fig. 1). This observation clearly indicates that armenite is orthorhombic and crystallizes in a primitive lattice. Systematic extinctions lead to the space-group Pnna ($\underline{a} = 13.874[2], \underline{b} = 18.661[2], \underline{c} = 10.697[1]$ Å). The structure of armenite was then solved in this space group by routine X-ray techniques described in a forth coming paper. A projection along a of the refined structure is shown in figure 2. Three different types of tetrahedra could be distinguished on the basis of T(Si,Al)-O distances. Of six symmetry distinct ring tetrahedra (T1 type), two show T-O distances characteristic of pure Si occupation whereas the remaining four indicate a Si/Al ratio of 1/1. Three tetrahedra (T2 type) connect the double-ring units to a tetrahedral framework. Two of the latter type also show a Si/Al ratio of 1/1 and one is occupied by Al. The (Si,Al) distribution is in agreement with a facecentered orthorhombic lattice. The observed primitive lattice is caused by the arrangement of H₂O molecules. In armenite, Ca is in seven-fold coordination where six oxygens of the tetrahedral framework form a distorted octahedron and one additional H₂O molecule completes the coordination. This H₂O molecule is either above or below the octahedron (in projections parallel to a, Pnna setting).

The refined structure (Fig. 2) shows a correlation between (Si,Al) ordering and H₂O arrangement. T2 type tetrahedra connect edge-sharing CaO₇ units with the water molecule on the same height along <u>a</u> and have a Si/Al ratio of 1/1. The tetrahedron which links two CaO₇ units with H₂O in opposite directions along <u>a</u>, has longer T–O distances and is indicative of pure Al occupation. This Al tetrahedron is also linked to four Si tetrahedra in the double-ring units.

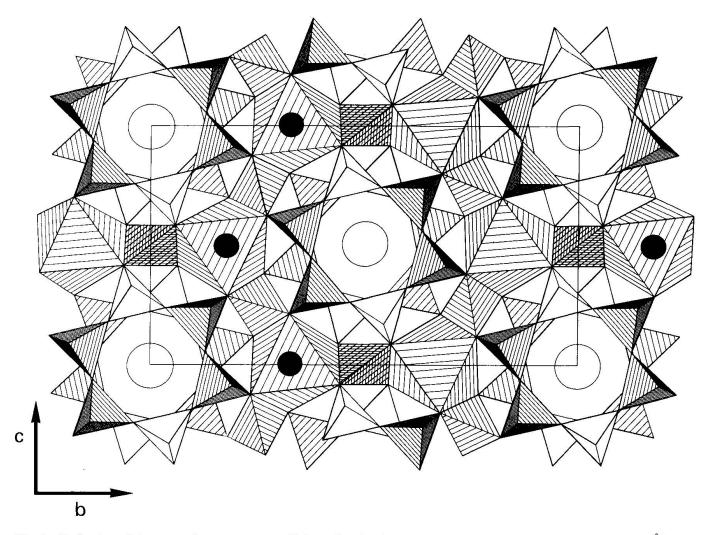


Fig. 2 Projection of the armenite structure parallel to \underline{a} (setting for space group Pnna: $\underline{a} = 13.8$, $\underline{b} = 18.6$, $\underline{c} = 10.7$ Å). In this projection, double-rings appear as single rings. Open tetrahedra are occupied by Si, hatched tetrahedra have Si/Al = 1/1, and cross-hatched tetrahedra are centered by Al. Ba (open circles) is shown between two adjacent double-ring units. Ca is in seven-fold coordination where six oxygens of the tetrahedral framework form a distorted octahedron and a H₂O molecule (filled circle) completes the coordination. In cases where no H₂O molecules are drawn the corresponding H₂O position is below the CaO₆ octahedron.

Previous heating experiments by NEUMANN (1941), Pouliot et al. (1984) and Balassone et al. (1989) showed that heating to 1000 °C expels H₂O from armenite. Our experiments on armenite single-crystals that were heated for 72 h at 1000 °C indicate that the superstructure reflections (characteristic of the primitive lattice) disappear and the crystals become metrically hexagonal. However, based on the experience of (Si,Al) orderdisorder in cordierite (e.g. Schreyer, 1966), it is not likely that these fairly moderate heating conditions are sufficient to cause complete (Si,Al) disorder within the double-ring (T1) and the linking tetrahedra (T2). Thus, armenites heated under these conditions are probably only pseudohexagonal and possess the same (Si,Al) arrangement as natural armenite. The corresponding space group symmetry is Amaa (in a setting corresponding to Pnna; standard setting is Cccm). The loss upon heating of H_2O leads to a pronounced shortening of \underline{b} and \underline{c} and an expansion of \underline{a} : a = 14.024(3), b = 18.383(3), c = 10.608(2) Å.

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