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Autor:	Ziegler, U.R.F.
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Preliminary results of geochemistry, Sm–Nd and Rb–Sr studies of post-Karoo carbonatite complexes in Southern Africa*

by U.R.F. Ziegler¹

Abstract

The preliminary results presented in this article show that the major and trace element composition of most of the analysed post-Karoo carbonatite complexes of Namibia (Kalkfeld), Zambia (Kaluwe, Chasweta), Malawi (Kangankunde) and Tanzania (Mbeya) is in good agreement with the expected composition for the respective carbonatite types. The chondrite normalized REE patterns of all the analysed carbonatites show a strong LREE/HREE enrichment and the absence of significant anomalies. K–Ar age determinations of the Eisenberg carbonatite (Kalkfeld complex) yielded ages between 153.6 and 172.8 Ma. Rb–Sr and Sm–Nd studies indicate depleted mantle reservoirs for the Kalkfeld, Kaluwe and Kangankunde carbonatites while the Mbeya complex seems to have a Rb-depleted, Nd-enriched reservoir. The study of the Chasweta is not conclusive with respect to its mantle source as the results show a bimodal sample distribution which either indicates different mantle sources or severe contamination/ mixing for the Chasweta magmas.

Keywords: Geochemistry, isotope geochemistry, carbonatites, post-Karoo complex, southern Africa.

1. Introduction

During the late Triassic the initial rifting of the South Atlantic triggered the separation of Africa from Antarctica and South America. Intracontinental stretching accompanying the opening of the South Atlantic and the South Indian Ocean led to a multitude of anorogenic magmatic events in southern Africa. During the early to late Jurassic the magmatism of the Karoo igneous province (ERLANK, 1984) was the characteristic and volumetrically most important type of magmatism in southern Africa. Old lines of weakness near the eastern end of Atlantic-related transform fault zones extending far inland (e.g. in Angola and Namibia) and intracontinental regions such as the Chilwa province in Malawi (e.g. WOOLLEY, 1991), the Rufunsa province in Zambia (BAILEY, 1966; TURNER, 1988) and the Mbeya area

in southern Tanzania (VAN STRAATEN, 1989) where lineaments and/or fault systems occur and partly transsect each other were the place of subsequent late Jurassic to early Cretaceous (ca. 165 to 100 Ma) ultra-alkalic and carbonatitic magmatism in southern Africa (KAMPUNZU and POPOFF, 1991). Within the last few years several studies on African carbonatites (e.g. BELL and BLENKINSOP, 1989; KINNAIRD and BOWDEN, 1991) have shown that the majority of African carbonatites have relatively smooth and steep REE patterns and that the source regions of the so far analysed African carbonatite complexes lie within the subcontinental mantle which must be inhomogeneous with respect to its Sr and Nd isotopic composition.

As no Rb–Sr, Sm–Nd, REE and age data exist on many of the major carbonatite complexes of southern Africa (e.g. the Kalkfeld complex in Namibia and the Rufunsa carbonatites of Zam-

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¹ Ittigenstrasse 11, CH-3063 Ittigen.

bia) while the isotope and trace element characteristics of the carbonatites of eastern Africa and the Republic of South Africa are partly well known there exists a major gap in the knowledge on the mechanisms and sources of the post-Karoo magmatism in these areas. The aim of the present study is thus to gain more information on these subjects on a regional scale by the analysis of the REE- and isotope systematics of various post-Karoo carbonatite complexes of southern Africa. This report summarises the first preliminary results of this ongoing study.

2. Sample description

The analysed samples cover all the major structural and tectonic settings in which Mesozoic carbonatites occur in southern Africa. They were collected in the following areas:

The Rufunsa province is located at the triple junction of the NE-SW Luangwa rift with the middle and lower Zambesi rifts (Fig. 1A). This province formed close to the notheastward extension of the (reactivated) Proterozoic Mwembeshi shear zone (DALY, 1988). Carbonatites occur in a group of four complexes emplaced into a sequence of Upper Karoo sandstones (BAILEY, 1966; TURNER and REX, in press). Detailed descriptions of the Rufunsa carbonatites which are not associated with any silicic magmatic rocks may be found in BAILEY (1960 and 1966), TURNER (1988). K-Ar age determinations of TURNER and REX (in press) yielded ages between 101 and 113 Ma for the Kaluwe complex. The samples which were analysed during this study derive from the northernmost Kaluwe and the southernmost Chasweta complexes. The Kaluwe samples Z19, Z20 and Z21 are dark grey, fine-grained sövites. The Chasweta samples Z1, Z4, Z12 are dark brown, fine-grained ferrocarbonatites while Z17 is a brownish, fine-grained sövite.

The Kalkfeld province (Fig. 1B) of central western Namibia is situated on the southwestnortheast trending lineament of the Damaraland province (PRINS, 1981) which represents an Atlantic related transform fault zone. The three carbonatite complexes of the Kalkfeld province (Kalkfeld, Ondurakorume, Osongombo) which have been described in detail by VERWOERD (1967) intruded the country rock of pan-African age amidst (on a regional scale) a large number of alkaline ring complexes. The analysed finegrained biotite sövite samples N15 and N33 were collected on the central Eisenberg plug of the Kalkfeld complex. The large (up to 3 cm in diameter), partly chloritized biotites (N36 I – N36 IV) used for the K-Ar age determinations of this study derive from the biotite sövites of the summit area of Eisenberg.

The Chilwa alkaline province (Fig. 1C) of southern Malawi is situated at the junction of the north-south trending East African Rift and the east-west trending Zambezi rift. Seventeen carbonatite centres of a wide compositional range are known from this area (GARSON, 1965). K-Ar age determinations of SNELLING (1965) and CAHEN and SNELLING (1966) on the Kangankunde, Tundulu and Chilwa island complexes yielded ages of 123 ± 6 Ma, 133 ± 7 Ma and 136 ± 7 Ma, respectively. The analysed samples M5, M7 and M8 are dark brown, fine-grained, monazite rich magnesiocarbonatites from Kangankunde Hill which is (besides some alnöitic dykes and some carbonated nephelinite) free of associated silicate rocks.

The Mbeya (Panda Hill) province is located in southwestern Tanzania south of the Tanzania craton along strike of the East African Rift (Fig. 1D). K–Ar age determinations of SNELLING (1965) and MILLER and BROWN (1963) yielded ages between 96 ± 9 Ma and 113 ± 6 Ma for the four Mesozoic carbonatites of southwestern Tanzania occurring in this area. The analysed samples EH12 and EH13 are creamy white to grey, finegrained sövites from the Panda Hill complex close to the city of Mbeya.

3. Methods

The major and REE elements of the samples N15, N33, Z1, Z4, Z12, Z17, Z19, Z20, Z21, M5 and M7 were determined at CRPG, Nancy, France using inductively coupled plasma atomic emission spectrometry while the REE composition of the samples EH12 and EH13 was determined by neutron activation at Bondar Clegg Ltd., Ottawa, Canada. CO₂ was coulometrically determined at CRPG in Nancy. All the isotopic measurements were carried out at the Laboratory for Isotope Geology of the University of Bern. The Rb-Sr concentrations and the Sr isotopic compositions were determined using the isotope dilution method as described by JAEGER (1979). The Sm and Nd concentrations as well as the Nd isotopic composition were determined according to the method described by Stoessel and Ziegler (1989) using a two column element separation procedure. The Rb analyses were carried out on an "Ion Instruments" solid source mass spectromenter. Elemental Sr, Sm, and Nd measurements were carried out on a "VG Sector" thermal ionisation mass spectrometer by single and triple filament



Fig. 1. Location map of the analysed complexes: A) Rufunsa area (Zambia) after BAILEY (1989); B) Kalkfeld province (Namibia) after KINNAIRD and BOWDEN (1991); C) Chilwa alkaline province (Malawi) after WOOLLEY (1991); D) Carbonatite areas of Tanzania after VAN STRAATEN (1989).

Comple	ex Kalki Namil	feld Dia		Chaswe Zambia	eta a		K Z	aluwe ambia		Kanga Mala	nkunde awi	Mbeya Tanza	ania
Sample	e N15	N33	Z 1	Z4	Z12	Z 17	Z19	220	Z21	M5	M7	EH12	EH13
Major	Elements	(wt%):											
SiO2 Al2O3 Fe2O3 MnO MgO CaO Na2O K2O TiO2 P2O5	0.79 0.13 3.72 0.91 0.21 51.82 0.07 tr tr 2.22	25.67 9.31 9.71 0.56 5.75 21.2 0.61 3.47 2.38 1.20	1.53 0.71 11.69 1.67 8.21 31.43 0.03 tr 0.11 1.18	3.11 1.45 11.94 4.25 38.37 0.03 0.11 0.07 1.29	1.63 0.78 11.35 2.02 5.05 37.26 0.04 tr 0.06 3.00	1.51 0.44 3.69 0.51 0.48 51.61 0.03 0.03 0.11 1.35	0.75 0.08 1.45 0.16 0.86 53.75 0.06 0.03 0.02 3.45	5.74 1.58 2.15 0.51 0.88 48.76 0.06 0.89 0.08 0.89 0.08	0.1 0.0 0.4 0.4 0.2 55.7 tr tr tr tr	4 5.99 7 1.56 8 4.76 8 1.64 712.03 521.66 0.05 0.14 0.02 1 2.84	4.99 1.41 10.64 2.51 11.71 26.87 0.02 tr 0.02 tr 0.02 1.08		
CO2	37.44	13.93	34.93	32.39	31.76	37.21	37.31	35.27	40.9	33.93	35.58		
Trace	Elements	(ppm):										51	
Ba Be Co Cr Cu Ga Nb Ni Rb Sc Sr Th V Y Zn Zr	434 2 19 13 49 54 189 16 1 26 8064 252 25 333 86 37	>5000 8 31 362 53 16 170 86 66 32 5844 301 148 86 127 275	>5000 7 42 27 155 35 >1000 19 <1 24 1830 216 151 141 510 105	>5000 10 69 26 57 32 460 157 6 26 1663 329 369 266 395 375	>5000 10 53 23 96 41 609 40 <1 42 1737 433 201 197 202 123	>5000 5 21 18 13 20 546 15 4 35 723 109 159 90 142 104	1119 1 19 26 19 962 14 2 16 5272 71 75 62 26 53	>5000 5 18 22 21 379 15 22 17 99 168 96 168 96 168 76	1278 2 11 19 19 362 10 1 20 748 55 104 74 45 32	3586 8 32 17 42 163 170 25 2 39 90258 822 43 80 >1000 153	>5000 3 22 21 31 79 23 13 1 28 27260 326 49 46 >1000 66	650 14 46 5 13278 4	1500 17 23 <1 4 3910 214
Rare e	arth elem	ents (p	pm):			*							
La Ce Nd Sm Eu Gd Dy Er	2967 5193 2897 335 75.8 212 113 47.6 23.9	183 285 180 35.7 9.7 23.7 13.2 7.1 6 1	2456 3743 1582 177 38.2 114 61.5 27.6	1825 2945 1267 161 41.7 127 77.5 40.3 29 7	2892 4195 1507 177 46.9 146 85.0 38.8 26 7	513 876 374 57.1 19.1 63.3 43.0 24.4	243 496 239 36.8 8.31 23.6 11.3 4.71	657 1045 439 63.1 20.5 60.8 45.8 23.1	402 724 294 43.9 10.0 28.3 15.0 6.4	14525 23261 8130 477 90 318 111 53 23	5625 9012 3129 216 53 144 65 58 20	612 1120 391 46.9 10	306 610 222 31.9 7
Lu Lu	3.9	1.1	2.9	5.5	5.0	3.0	0.33	3.0	0.4	23 5	4 	0.9	<0.2
			TOTAL STREET, ST		torrespective of the second of	the state of the second second	and the second sec	a a second to be the	and the second second		1 17 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5		

Tab. 1 Major, trace and rare earth element concentrations of the analysed samples.

modes, respectively. The 87 Sr/ 86 Sr ratios are normalized to 88 Sr/ 86 Sr = 0.1194 while the 143 Nd/ 144 Nd ratios are normalized to a 146 Nd/ 144 Nd ratio of 0.72190. Total blanks were < 1 ng for Sm, 0.01–0.1 ng for Nd and < 10 ng for Sr. K was analysed by flame photometry from an aliquot of the sample. The isotopic composition and the concentration of Ar were determined on a VG MM1200 solid source mass spectrometer following the extraction and purification in a Pyrex glass line and the addition of 99.997% pure 38 Ar spike.

4. Results and discussion

The results of the major, trace and rare earth element analyses are listed in table 1 while those of the isotopic analyses are compiled in tables 2-4.

The K-Ar ages of the analysed Eisenberg samples range between 153.6 and 172.8 Ma and thus confirm the classification of the Kalkfeld carbonatites as post-Karoo complexes.

Figure 2 shows a general classification of the analysed samples according to the propositions of WOOLLEY and KEMPE (1988). A comparison with the data in WOOLLEY and KEMPE (1988) shows, with the exception of N33, that the major and trace element geochemistry of the samples is (although certain trace elements are in cases slightly enriched compared to the literature data) generally in good agreement with the expected range for the respective carbonatite types. The abnormal composition of N33 is due to the presence of biotite, chlorite, vermiculite, augite and some

5							
Sample	Rb ppm	Sr ppm	87Sr/86Sr	+/-	87Rb/86Sr	+/-	Epsilon Sr(T)
Kalkfeld Namibia			е Т		÷		
N15 N33	0.57 65.53	8064 5844	0.703669 0.703918	10 12	0.000206 0.032432	0.000002 0.000324	-9.1 -6.6
Chasweta Zambia							
Z1 Z4 Z12 Z17	0.32 6.19 0.41 3.68	1830 1663 1737 723	0.707739 0.705736 0.706338 0.703836	9 7 11 8	0.000498 0.010769 0.000683 0.014702	0.000005 0.000108 0.000007 0.000147	47.9 19.2 28.0 -7.9
Kaluwe Zambia							
z19 z20 z21	2.18 22.89 1.08	5272 992 748	0.703286 0.704240 0.703747	9 11 8	0.001195 0.066747 0.004169	0.000012 0.000667 0.000042	-15.4 -3.3 -8.9
Kangankun Malawi	de			12			
M5 M7 M8	1.82 0.81 0.59	90258 27260 39367	0.703016 0.703067 0.703033	5 7 9	0.000059 0.000086 0.000043	0.000001 0.000001 0.000001	-19.0 -18.3 -18.7
Mbeya Tanzania							
EH12 EH13	5.69 0.61	13279 3910	0.703876 0.703407	7 8	0.001238 0.000449	0.000012 0.000004	-7.2 -13.9

Tab. 2 Tabulation of the Rb and Sr results. Constants used are: ⁸⁷Rb decay constant = $1.42 \cdot {}^{-11}a^{-1}$ (STEIGER and JAEGER, 1977), ⁸⁵Rb/⁸⁷Rb = 2.59265, ⁸⁶Sr/⁸⁸Sr = 0.1194, ⁸⁴Sr/⁸⁶Sr = 0.056584 (STEIGER and JAEGER, 1977), ⁸⁷Sr/⁸⁶Sr_{Uniform Reservoir UR} = 0.7045 (e.g. DE PAOLO, 1988), ⁸⁷Rb/⁸⁶Sr UR = 0.0827 (e.g. DE PAOLO, 1988).

K-feldspar besides carbonates and limonite. Figure 3 shows the results of the REE analyses in chondrite normalized abundance patterns. All the REE abundance patterns are very steep and smooth due to very high light REE abundances and the absence of pronounced anomalies both of which are typical features of carbonatites. The ϵ Nd_(T) values calculated for the respective intrusion ages of the analysed complexes from the measured ¹⁴³Nd/¹⁴⁴Nd data range between -5.3 and +5.5. The Kangankunde, Kaluwe and Eisenberg samples all have positive ϵ Nd_(T) values while those of the Mbeya and Chasweta samples are at least partially negative. With the exception of Chasweta all the ϵ Sr_(T) values of the Chasweta samples are between -7.9 and +47.9.

Figure 4 shows the results of the isotopic measurements in an $\epsilon Nd_{(T)}$ v. $\epsilon Sr_{(T)}$ diagram. All the Kaluwe, Eisenberg and Kangankunde samples as well as one Chasweta data point scatter within the upper left (depleted) quadrant. This indicates that



Fig. 2 CaO-MgO-Total Iron Oxide + MnO plot of the analysed carbonatites (same symbols as in Fig. 4).

Tab. 3 Tabulation of the Sm and Nd results. Constants used are: ¹⁴⁷Sm decay constant = 0.00654 b.y.⁻¹ (LUGMAIR and MARTI, 1978), ¹⁴³Nd/¹⁴⁴Nd_{Chondritic Uniform Reservoir CHUR (t = 0)} = 0.51264, ¹⁴³Nd/¹⁴⁴Nd_{Depleted Mantle DM(0)} = 0.51316, ¹⁴⁷Sm/¹⁴⁴Nd_{CHUR} = 0.1967, ¹⁴⁷Sm/¹⁴⁴Nd DM = 0.2136.

				1993 - 19 - 199 2 - 199			
Sample	Sm ppm	Nd ppm	143Nd/144Nd	+/-	147Sm/144	1Nd +/-	Epsilon Nd(T)
Kalkfeld Namibia					đ		5. 5.
N15 N33	334.65 35.74	2897.8 180.06	0.512575 0.512580	2 3	0.0698 0.1200	0.0007 0.0012	1.3 0.4
Chasweta Zambia	<i>t</i> i	,		<u>.</u> .			
Z1 Z4	177.00	1582.1 1267.0	0.512442	2	0.0676	0.0007	-2.1 -2.8
Z12 Z17	176.95 57.12	1506.9 374.48	0.512280 0.512849	5 16	0.0710 0.0922	0.0007 0.0009	-5.3 5.5
Kaluwe Zambia							
Z19 Z20 Z21	36.83 63.07 43.95	238.88 439.67 293.51	0.512737 0.512694 0.512750	18 3 3	0.0932 0.0867 0.0905	0.0009 0.0009 0.0009	3.3 2.6 3.6
Kangankun Malawi	de						
м5 м7 м8	477.18 216.33 189.05	8129.6 3129.3 2950.3	0.512679 0.512694 0.512669	2 3 2	0.0355 0.0418 0.0387	0.0004 0.0004 0.0004	3.3 3.5 3.0
Mbeya Tanzania		c.					
EH12 EH13	46.88 31.85	390.76 222.00	0.512511 0.512506	3 5	0.072523 0.086734	0.0007 0.0009	-1.0 -1.3

Tab. 4 Tabulation of the K–Ar results. Constants used are: 40 K = 5.543 • 10⁻¹⁰a⁻¹ (total 40 K decay constant), 40 K^e = 0.581 • 10⁻¹⁰a⁻¹ (STEIGER and JAEGER, 1977), 40 Ar/ 36 Ar_{atm} = 295.5, 38 Ar/ 36 Ar_{atm} = 0.1869 (STEIGER and JAEGER, 1977).

Sample		K(wt%)	40Ar/36Ar	40K/36Ar	Age (Ma)
N36	т	1 41	1243 8	94898	164 3
N36	ī	2.44	1069.5	78984	161.3
N36	III	3.37	3620.4	315562	172.8
N36	IV	2.58	6144.8	627735	153.6

these carbonatites derive from reservoirs depleted in Large Ion Lithophile Elements (LILE) relative to bulk Earth while the Mbeya samples seem to derive from a mantle source with lower Sm/Nd and Rb/Sr ratios relative to the bulk Earth. Three of the four Chasweta samples plot within the enriched quadrant. A comparison with literature data taken from a compilation of BELL and BLENKINSOP (1989) shows that the majority of the samples plots in the areas of known carbonatite compositions although some of them do not follow the East African carbonatite line of BELL

and BLENKINSOP (1989). The high ⁸⁷Sr/⁸⁶Sr and low ¹⁴³Nd/¹⁴⁴Nd ratios of the three enriched Chasweta samples are somewhat puzzling as they plot far away from known carbonatite compositions. The fact that the fourth Chasweta sample (Z17) as well as all the Kaluwe samples do not show such aberrant isotopic compositions raises even more questions on the genesis of these Rufunsa carbonatites. The only conclusions that can be drawn from the present data are the following: The Kaluwe samples as well as Z17 seem to have their origins in a mantle reservoir depleted in LILE and their isotopic composition remained more or less unchanged during formation, ascent and emplacement of the carbonatite magmas. If the sövitic Z17 and the three enriched ferrocarbonatitic Chasweta samples which show a certain tendency towards Group II kimberlites (data from GURNEY et al., 1991) do not have the same origins then the aberrant samples must either derive from a comparatively enriched mantle source or they must have had a strong crustal component in their



Fig. 3 Chondrite normalized REE plot for the analysed carbonatites.



Fig. 4 Plot of $\varepsilon Nd_{(T)}$ versus $\varepsilon Sr_{(T)}$ for the analysed samples. Data for comparison are taken from the compilation of Bell and BLENKINSOP (1989).

source area while Z17 derives from a depleted source. If, however, they have common origins then certain parts of the Chasweta magma must have been subject to mixing processes including the assimilation of continental crust material. The first interpretation is supported up to a certain degree by the considerations of BAILEY (1989) who identified magnesio-chromite-bearing pyroclasts in part of the Chasweta carbonatites which thus indicate a certain kimberlite affinity of these pyroclastic carbonatites. As magnesio-chromite is only present in the pyroclastic carbonatites and is absent in the coarse grained intrusive carbonatites BAILEY (1989) also stated that the Rufunsa carbonatites must derive from various origins; a conclusion which would also support the first interpretation. However, only further detailed field and laboratory work can help to find a solution to this Rufunsa problem.

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