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Johninnesite, $\text{Na}_2(\text{Mn}^{2+})_9(\text{Mg,Mn})_7(\text{AsO}_4)_2(\text{Si}_6\text{O}_{17})_2(\text{OH})_8$: a new occurrence in Val Ferrera (Graubünden, Switzerland)

by Joël Brugger^{1,2} and Peter Berlepsch¹

Abstract

The rare mineral johninnesite, $\text{Na}_2(\text{Mn}^{2+})_9(\text{Mg,Mn})_7(\text{AsO}_4)_2(\text{Si}_6\text{O}_{17})_2(\text{OH})_8$, has been discovered within a metamorphic Fe-Mn-deposit in Triassic marbles of the Schams nappes in Val Ferrera (Graubünden, Switzerland). To our knowledge, this is the second occurrence of this mineral, which was first described from Fe-Mn-ore lenses at the Kombat mine (Namibia). Johninnesite from Val Ferrera differs from the type material by its Fe and V contents, and by higher MgO contents which are correlated to lower MnO contents. IR and Raman spectra for this species are reported for the first time. Johninnesite is a product of the metamorphism of syn-genetic exhalative deposits, and the reason for the rarity can be ascribed to the geochemistry of the protolith. In particular, if the protolith is not depleted in aluminium, the widespread silico-arsenate ardennite will form instead of johninnesite.

Keywords: johninnesite, chemical analyses, arsenic, IR and Raman spectra, Val Ferrera, Switzerland.

Introduction

The rare mineral johninnesite has been described by DUNN et al. (1986) from the Kombat Mine (Namibia). This mine is located in dolostones of the Hüttenberg formation, the youngest formation of the Upper Proterozoic Otavi Group (Damara sequence, South-West Africa/Namibia; INNES and CHAPLIN, 1986). Johninnesite occurs in Fe-Mn-ore lenses embedded in feldspathic sandstones deposited near, or at the top of the Hüttenberg formation and subsequently transposed into the dolostones along steep folds. The Fe-Mn ores are layered and consist of magnetite, hausmannite, hematite, barite, calcite, tephroite, alleghanyite, pyrochroite. Their ore texture, mineralogy and geochemistry suggest that they are the metamorphosed equivalents of volcanic exhalative deposits.

The Hüttenberg formation underwent subgreenschist facies metamorphism. At the Kombat mine, ubiquitous rock alteration is linked to the deposition of various Cu, Pb, and Ag ores. The

temperature of the main mineralizing event has been measured from fluid inclusions as 200–280 °C (INNES and CHAPLIN, 1986).

The Fe-Mn ores of the Kombat mine are famous because they furnished a number of rare and unusual minerals, including Mn-bearing silicates such as leucophoenicite, glaucochroite, cuspidine, and ribbeite (type locality, PEACOR et al., 1987), vanadates such as kombatite (type locality, ROUSE et al., 1986), and arsenates such as sahlinite and johninnesite.

The structure of johninnesite was solved by GRICE and DUNN (1994), leading to the ideal structural formula $\text{Na}_2\text{Mn}_9(\text{Mg,Mn})_7(\text{AsO}_4)_2(\text{Si}_6\text{O}_{17})_2(\text{OH})_8$. The mineral is triclinic with space group P1. The johninnesite structure consists of sheets parallel to (100). GRICE and DUNN (1994) distinguish three distinct kinds of polyhedral sheets: (1) the tetrahedral layer (T) shares $[\text{Si}_6\text{O}_{17}]^{10-}$ rings and isolated $[\text{AsO}_4]^{3-}$ groups; (2) an edge sharing octahedral layer (O) with brucite-like geometry, 5/12 positions being occupied by Mn, 5/12 by (Mg,Mn) and the remaining two be-

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ing empty; (3) another octahedral layer (O') sharing Mn and Na polyhedra. The stacking order is TOTO'.

During the systematic study of the small Fe and Fe–Mn deposits of the Val Ferrera (Graubünden, Switzerland), johnninesite was found in abundance. As far as we know, Val Ferrera is the second locality for the mineral. This paper describes briefly the mineral assemblage and the geological context of this occurrence, discusses new chemical data on johnninesite (including johnninesite from the type locality), and gives some new spectroscopic data for this rare mineral species.

Occurrence

Numerous small Fe and Fe–Mn deposits (STUCKY, 1960) occur in the Triassic dolomite and calcite marbles of Val Ferrera (Graubünden, Switzerland), which belong to the para-autochthonous sedimentary cover of the Suretta nappe, to the Starlera nappe, and to the Schams nappes

(BAUDIN et al., 1995; SCHMID et al., in press). A syn-genetic exhalative origin has been proposed for these deposits (BRUGGER, 1996), which subsequently underwent a complex Tertiary metamorphism climaxing under greenschist to blueschist facies conditions (MARQUER et al., in press; SCHREURS, 1991).

So far, johnninesite has only been identified in a restricted area (about 10 m) of the Schmorrasgrat deposit in the Schams nappes. The host ore consists of hematite, jacobsite, quartz, dolomite, kutnohorite, and amphibole (magnesian riebeckite), which is often brecciated so that quartz and dolomite grains up to 2 cm are embedded in a black matrix of the ore minerals. Johnninesite may be the main component, with minor quartz and carbonate, in discordant veins up to 5 cm in width, or it may occur as an accessory in quartz, carbonate, manganberzeliite-palenzonaite veins. Rare associated minerals in both kinds of veins are walkkildellite, pink manganian biotite and fluoromérite (BRUGGER et al., 1997). Three johnninesite morphologies can be distinguished: (a) fi-



Fig. 1 Microphotograph of a johnninesite veinlet. j = johnninesite, r = roméite, d = dolomite (belongs to the veinlet), g = embedding ore, consisting mostly of dolomite, hematite, jacobsite, and manganian muscovite. Transmitted light, field of view 1.25×0.92 mm.

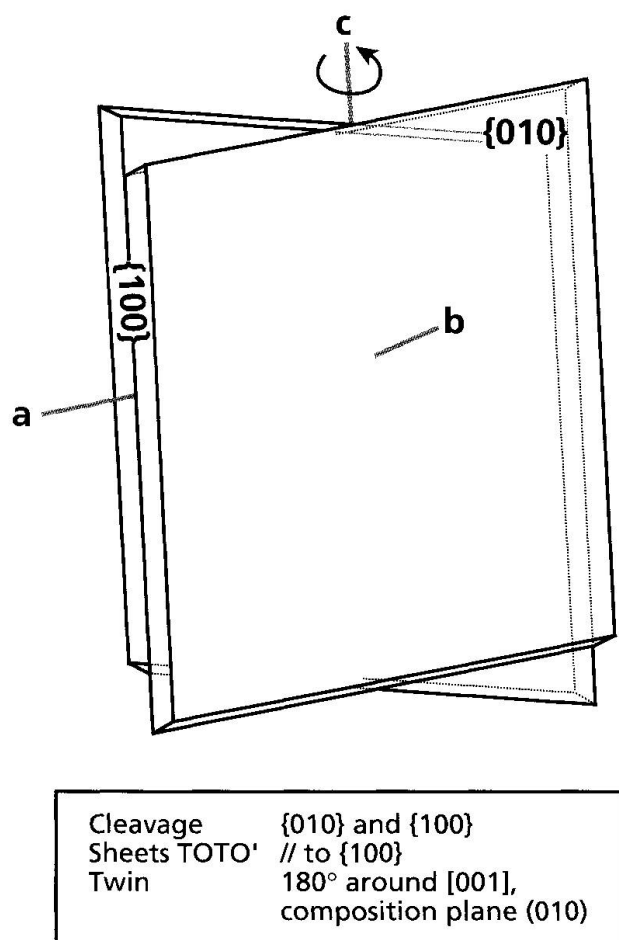


Fig. 2 Schematic view of a johninnesite crystal with a morphology of type (b). Polysynthetic twinning occurs by a 180° rotation around [001]. The dark/light gray bands represent the stacking of the polyhedral sheets which build the johninnesite structure.

brous crystals up to 4 cm in length, elongated perpendicular to the vein walls, which with quartz + carbonate grew synchronously to the opening of the fracture (Fig. 1); (b) rarely, similar veinlets contain prismatic crystals (about 2 mm × 1 mm × 0.3 mm; Fig. 2); and (c) small micaceous crystals (< 1 mm) which fill thin fractures over several cm².

Crystallography

Johninnesite from Val Ferrera originally was thought to be a yellow amphibole. A routine X-ray powder diagram was obtained with a Gandolfi camera, and a database check with the four strongest lines from the powder pattern unequivocally led to the triclinic mineral johninnesite. Powder data for johninnesite from the Kombat mine and from Val Ferrera are listed in table 1.

Tab. 1 Powder data for johninnesite from Val Ferrera and from the type locality.

| Val Ferrera ¹ | | | Kombat Mine ² | |
|--------------------------|------------------|-------------------|--------------------------|------------------|
| d_{obs} | I_{obs} | d_{calc} | d_{obs} | I_{obs} |
| 10.08 | 20 | 9.83 | | |
| 9.67 | 30 | 9.71 | 9.8 | 60 |
| | | | 7.8 | 2 |
| | | | 7.3 | 5 |
| 5.93 | 20 | 5.928 | 5.99 | 40 |
| 5.90 | 25 | | | |
| 5.193 | < 10 | 5.197 | 5.26 | 10 |
| 4.859 | < 10 | 4.855 | 4.91 | 10 |
| | | | 4.39 | 10 |
| | | | 4.24 | 2 |
| 4.046 | < 10 | 4.042 | 4.06 | 30 |
| 3.862 | 25 | 3.863 | 3.89 | 30 |
| 3.659 | 10 | 3.666 | 3.67 | 30 |
| 3.451 | 10 | 3.451 | 3.48 | 30 |
| 3.351 | 20 | 3.358 | 3.38 | 40 |
| 3.257 | 20 | 3.276 | | |
| 3.213 | 50 | 3.237 | 3.23 | 60 |
| 3.037 | 10 | 3.037 | 3.05 | 10 |
| 2.945 | 20 | 2.955 | 2.965 | 20 |
| 2.853 | 20 | 2.852 | 2.892 | 20 |
| 2.829 | 20 | 2.828 | 2.842 | 20 |
| 2.667 | 100 | 2.662 | 2.676 | 100 |
| 2.620 | < 10 | 2.621 | 2.63 | 2 |
| 2.561 | 20 | 2.561 | 2.571 | 10 |
| 2.474 | 30 | 2.462 | 2.479 | 40 |
| 2.429 | 20 | 2.429 | 2.436 | 30 |
| 2.352 | 30 | 2.351 | 2.354 | 20 |
| | | | 2.288 | 10 |
| 2.222 | 30 | 2.222 | 2.228 | 20 |
| 2.182 | 20 | 2.182 | 2.19 | 10 |
| 2.139 | 20 | 2.139 | 2.147 | 5 |
| 2.035 | 20 | 2.035 | 2.044 | 20 |
| 1.859 | 30 | 1.859 | 1.866 | 30 |
| 1.791 | 20 | 1.792 | 1.796 | 20 |
| 1.723 | 10 | 1.723 | 1.726 | 10 |
| 1.695 | 10 | 1.695 | 1.701 | 10 |
| 1.623 | 10 | 1.623 | 1.623 | 2 |
| 1.604 | 50 | 1.604 | 1.606 | 30 |
| 1.538 | 50 | 1.538 | 1.539 | 40 |
| | | | 1.498 | 2 |
| 1.438 | < 10 | 1.439 | 1.441 | 10 |
| | | | 1.427 | 5 |
| 1.398 | < 10 | 1.398 | 1.402 | 2 |
| 1.380 | < 10 | 1.381 | 1.386 | 2 |
| 1.362 | 30 | 1.362 | 1.365 | 5 |

¹ G683, Schmorrasgrat; Gandolfi camera 114.6 mm diameter, Mn-filtered FeK_α radiation ($\lambda = 1.93728 \text{ \AA}$); I_{obs} : visually estimated intensities; d_{calc} : theoretical d-values for the cell obtained by least-squares refinement of G683 (program Appleman, APPLEMAN and EVANS, 1973).

² DUNN et al. (1986): 4 d_{obs} with $I_{\text{obs}} = 1$ are not listed.

The unit cell parameters (Tab. 2) were derived by least-squares refinement using the powder data and compared to the data reported by DUNN

Tab. 2 Cell parameters for johninnesite from Val Ferrera and from the type locality.

| | Val Ferrera | Kombat mine |
|--------------|-------------|-------------|
| a [Å] | 10.48(1) | 10.44(2) |
| b [Å] | 11.01(1) | 11.064(6) |
| c [Å] | 9.58(1) | 9.62(1) |
| α [°] | 107.02(2) | 107.43(7) |
| β [°] | 82.54(2) | 82.7(1) |
| γ [°] | 112.03(1) | 111.61(1) |

et al. (1986). Taking into account the variation in chemical composition, the refined unit cells are in good agreement.

The same twinning as reported by DUNN et al. (1986) and GRICE and DUNN (1994) was observed on a precession film showing the (100) plane of the reciprocal lattice of johninnesite from Val Ferrera (Fig. 2). This film also displayed weak reflections which could be related to a super structure.

Chemistry

Chemical analyses were performed on a JEOL JXA-8600 electron microprobe operated at 15 kV and 10 nA with a focused electron beam scanning a surface of about 20 μm^2 . The following stan-

dards, lines, and analysator crystals were used: adamite: As L_{α} /TAP; V_2O_5 : V K_{α} /LiF; albite: Si K_{α} /TAP; spessartine: Mn K_{α} /LiF; graffonite: Fe K_{α} /LiF; olivine: Mg K_{α} /TAP; albite: Na K_{α} /TAP; wollastonite: Ca K_{α} /TAP; celestite: Sr K_{α} /PET; galena: Pb L_{α} /PET.

This analytical procedure yields satisfactory results for minerals such as amphiboles, medaite or ardennite. However, the johninnesite analyses (Tab. 3) are characterized, relative to the theoretical formula, by an excess silicon content which correlates mainly with deficiencies of the (Mg+Mn) values. The use of different standards (e.g., marokite for Mn, sillimanite or jadeite for Si) could not eliminate this non-stoichiometry.

The comparison of the analyses of johninnesite from the Kombat mine and from Val Ferrera obtained under similar analytical conditions shows that:

(1) At Val Ferrera it contains small but significant amounts of Fe.

(2) It also has a significant V content, which is negatively correlated to the As content; V therefore enters the johninnesite structure along the As_{-1}V exchange vector (Fig. 3a).

(3) Johninnesite from Val Ferrera usually has a higher MgO content which correlates with a lower MnO content (Fig. 3b).

Tab. 3 Chemical data for johninnesite (all values in weight-%).

| | σ^1 | A | B | range | C | range |
|-------------------------|------------|-------------------|--------|--------------|-------|--------------|
| SiO_2 | 0.21 | 35.5 | 35.68 | 35.35–36.06 | 36.34 | 35.66– 37.35 |
| As_2O_5 | 0.18 | 10.6 | 9.43 | 9.01– 9.87 | 8.59 | 6.63– 9.81 |
| V_2O_5 | 0.10 | | 0.13 | < 0.10– 0.18 | 0.81 | 0.30– 1.34 |
| MnO | 0.41 | 40.7 | 39.04 | 38.25–39.97 | 37.91 | 35.97– 39.95 |
| FeO | 0.05 | 0.1 | < 0.05 | < 0.05– 0.13 | 0.20 | < 0.05– 1.86 |
| MgO | 0.10 | 8.2 | 7.12 | 6.92– 7.24 | 8.43 | 7.26– 9.30 |
| CaO | 0.05 | | < 0.05 | – | 0.13 | < 0.05– 0.68 |
| Na_2O | 0.07 | 3.1 | 2.94 | 2.85– 3.02 | 2.84 | 2.64– 2.99 |
| H_2O^2 | | 3.51 ³ | 3.38 | 3.39– 3.43 | 3.47 | 3.42– 3.53 |
| Sum | | 101.71 | 97.75 | 96.95–98.18 | 98.72 | 97.72–100.29 |

PbO: up to 0.15(5) wt% in Val Ferrera; < 0.05 wt% in the Kombat mine.

Ideal formula (GRICE and DUNN, 1994): $\text{Na}_2\text{Mn}_9(\text{Mg},\text{Mn})_7(\text{AsO}_4)_2(\text{Si}_6\text{O}_{17})_2(\text{OH})_8$

Empirical formulas:

A: $\text{Na}_{2.05}\text{Mn}_{9.00}(\text{Mg}_{4.18}\text{Mn}_{2.78}\text{Fe}_{0.03})_{6.99}\text{As}_{1.89}\text{Si}_{12.13}\text{O}_{42}(\text{OH})_8$

B: $\text{Na}_{2.01}\text{Mn}_{9.00}(\text{Mg}_{3.74}\text{Mn}_{2.66}\text{Fe}_{0.01})_{6.41}(\text{As}_{1.74}\text{V}_{0.03})_{1.77}\text{Si}_{12.58}\text{O}_{42}(\text{OH})_8$

C: $\text{Na}_{1.90}\text{Mn}_{9.00}(\text{Mg}_{4.35}\text{Mn}_{2.11}\text{Fe}_{0.06})_{6.52}(\text{As}_{1.55}\text{V}_{0.19})_{1.74}\text{Si}_{12.57}\text{O}_{42}(\text{OH})_8$

A Type locality, DUNN et al. (1986).

B Type locality, this paper; mean of 6 analyses.

C Schmorrasgrat, Graubünden, Switzerland; mean of 16 analyses.

¹ error on counting statistics; for trace elements, this corresponds to the detection limit.

² calculated with respect to the ideal formula.

Normalization on 50 oxygens per formula unit (including 8 hydroxyl groups).

³ the value measured by DUNN et al. (1986) was 2.6 wt% H_2O .

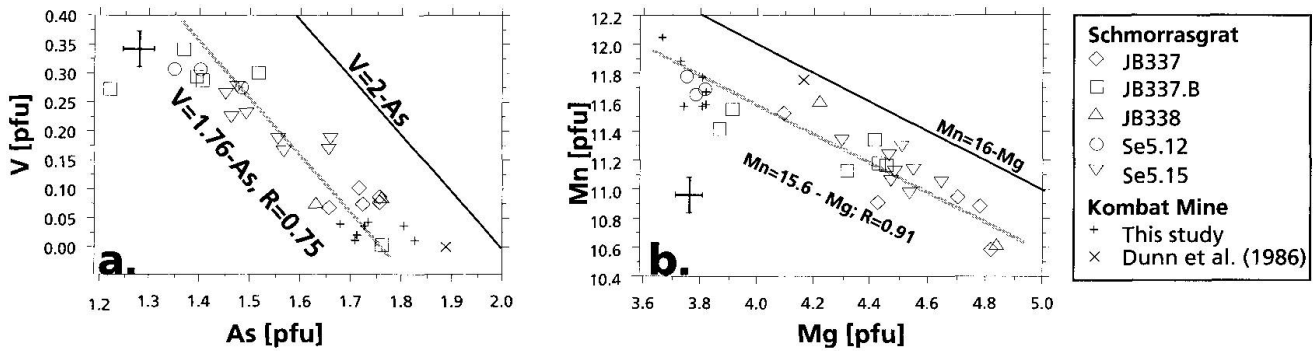


Fig. 3 Chemical variation of johninnesite (normalized EMP data). 1- σ errors from the counting statistics are plotted.

Infrared and Raman spectroscopy

Fourier transform infrared (FTIR) spectra were obtained on a PHILIPS PU9800 spectrometer using KBr pills (5 mm \varnothing , 60 mg KBr, ~ 0.1 mm³ johninnesite; 13 mm \varnothing , 220 mg KBr, ~ 0.5 mm³ johninnesite). The spectra of johninnesite from the type locality and from Val Ferrera mainly differ by an absorption peak at 1437 cm⁻¹ (Val Ferrera), which becomes very weak on the spectrum obtained on the co-type material (Fig. 4). Note that the IR-spectra of johninnesite and of saneroite [Na₂Mn₁₀VSi₁₁O₃₄(OH)₄] from the Gambatesa mine (Northern Italy) are very similar.

Raman spectra were collected on a RENISHAW spectrometer using the green line of an argon laser ($\lambda = 514.5$ nm). Both spectra are very similar (Fig. 5).

Discussion

Johninnesite is one of those minerals that occur in relatively large quantities but only in a very limited number of deposits world-wide. So far it has been described only from two carbonate-hosted Fe-Mn-deposits of syngenetic exhalative origin, which underwent regional metamorphism (Val Ferrera) or extensive hydrothermal alteration (Kombat mine).

We suggest that the geochemistry of the protolith is the most important limiting parameter responsible for the rarity of this mineral. A large positive As anomaly is characteristic for hydrothermal exhalative Mn-deposits (MARCHIG et al., 1982). Most of the deposits of this type are, however, related to hydrothermal activity at oceanic ridges (e.g., BONATTI, 1975; PETERS et al., 1980), and are embedded in deep sea sediments such as radiolarites. In contrast to the carbonate-hosted deposits in which johninnesite occurs,

oceanic deposits display higher Al contents (related to a higher detrital component); this fact may explain why the Al-mineral ardenite (Tab. 4), which has not been identified in the carbonate-hosted deposits of Val Ferrera and Kombat, is one of the most common silico-arsenates in metamorphic Mn-deposits. It has been recognized in many radiolarite-hosted Mn-deposits in the Alps, e.g., from Graubünden (Falotta: NIEVERGELT, personal communication [GEIGER and CABALZAR, 1989]; Parsettens: GEIGER and CABALZAR, 1989; Fur-

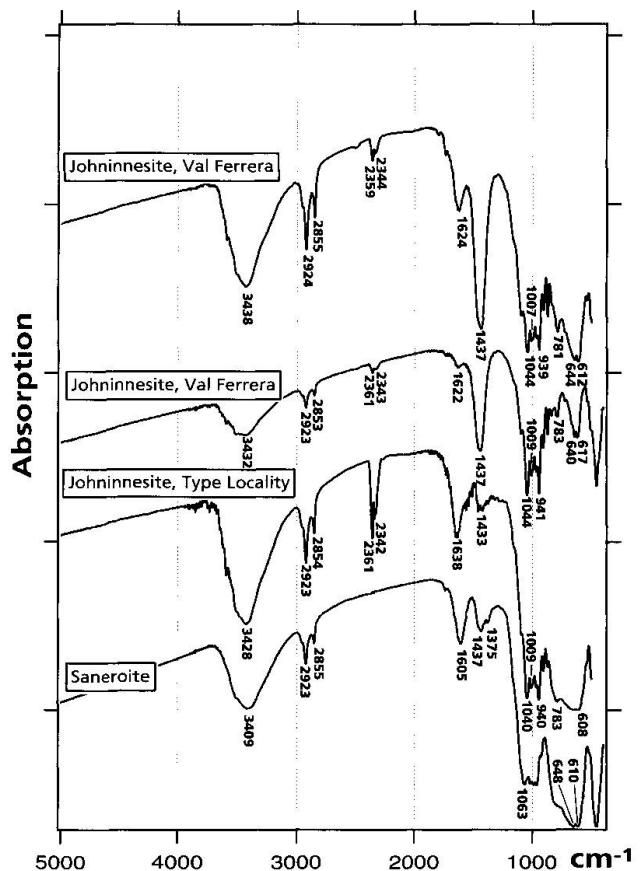


Fig. 4 IR-spectra of johninnesite from the type locality (Kombat mine), and from Schmorrasgrat, as well as of saneroite from the type locality (Val Graveglia, Italy).

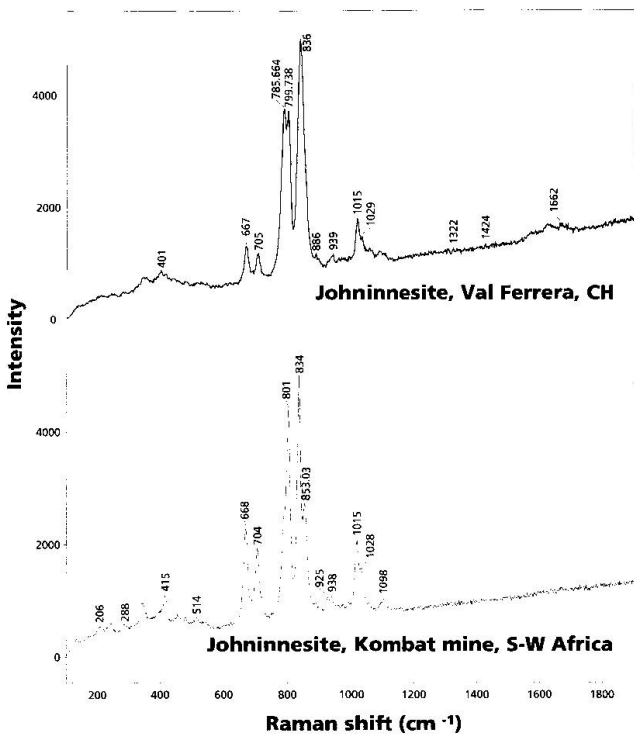


Fig. 5 Raman spectra of johnninesite from the type locality (Kombat mine), and from Schmorrasgrat.

tschellas: DE CAPITANI, 1983; Piz Cam: NIEVERGELT, personal communication [GEIGER and CABALZAR, 1989]; Northern Italy (Praborna, Val d'Aoste: MARTIN-VERNIZZI, 1984; Vedretta di Scerscen inferiore, Val Malenco: BEDOGNÉ et al., 1993), France (Haute Maurienne: CHOPIN, 1978), and in Greece (Southern Evia: REINECKE, 1986). The carbonate-hosted Mn-deposits are sometimes rich in vanadium. At Fianel, in another Fe-Mn deposit in the Val Ferrera (BRUGGER and BERLEPSCH, 1996), medaite is a very common mineral, and it seems that there the V/As ratio largely prevented the appearance of silico-arsenates. Johnninesite is also the only silico-arsenate which contains Na as a main constituent. This can be related to the geochemistry of the protolith, but also to the nature of the fluids from which johnninesite crystallized.

This study also draws attention to the difficulties in analyzing these rare minerals, which also have not received extensive crystal-chemical study. In particular, the discrepancy between the EMP analyses from this study and those of DUNN et al. (1986) is quite appreciable (Tab. 3). The non-stoichiometric analyses we obtained may be an artifact induced by the choice of the EMP standards. However, this is unlikely because the same calibration furnished good results for several other arseno-silicates. If the non-stoichiometry is real, it could be related to the presence of polysomatic

Tab. 4 The known manganese arseno-silicates and some manganese vanado-silicates.

| Mineral | Chemical formula |
|--------------|--|
| Ardennite | $(\text{Mn}^{2+}, \text{Ca})_4(\text{Al}, \text{Mn}^{3+})_5\text{Mg}(\text{As}, \text{V})\text{Si}_4\text{O}_{18}(\text{OH}, \text{O})_{10}$ |
| Johnninesite | $\text{Na}_2(\text{Mn}^{2+})_9(\text{Mg}, \text{Mn})_7(\text{AsO}_4)_2(\text{Si}_6\text{O}_{17})_2(\text{OH})_8$ |
| Nellerite | $(\text{Mn}, \text{Fe})_{16}\text{As}_3\text{Si}_{12}\text{O}_{36}(\text{OH})_{17}$ |
| Schallerite | $(\text{Mn}, \text{Fe})_{16}\text{As}_3\text{Si}_{12}\text{O}_{36}(\text{OH})_{17}$ |
| Tiragalloite | $\text{Mn}_4\text{AsSi}_3\text{O}_{12}(\text{OH})$ |

disorder. To test this, a preliminary transmission electron microscope (TEM) investigation has been performed by Allan Pring (personal communication). Although interpretation of these data is complicated by the triclinic geometry and the fine twinning of the mineral, they do suggest that samples from Val Ferrera are subject to some kind of polysomatic or polytypic disorder. This aspect of the work is still in progress.

Acknowledgement

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References

- APPLEMAN, D.E. and EVANS, H.T.J. (1973): Job 9214: indexing and least squares refinement of powder diffraction data. Natl. Tech. Inf. Serv. U.S. Dep. Commerce, Springfield, Virginia, Document PB-216 188.
- BAUDIN, T., MARQUER, D., BARFETY, J.-C., KERCHKOVÉ, C. and PERSOZ, F. (1995): A new stratigraphical interpretation of the Mesozoic cover of the Tambo and Suretta nappes: evidence for early thin-skinned tectonics (Swiss Central Alps). *Comptes rendus de l'Académie des Sciences, Paris, France* t. 321, série IIa, 401–408.
- BEDOGNÉ, F., MONTRASIO, A. and SCIESA, E. (1993): I Minerali della provincia di Sondrio, Valmalenco. Bettini, Sondrio-Italy, 275 pp.
- BONATTI, E. (1975): Metallogensis at oceanic spreading centers. *Ann. Rev. Earth Plan. Sci.*, 3, 401–431.
- BRUGGER, J. (1996): The Fe, Mn, (V, Sb, As, Be, W) deposits of Val Ferrera (Graubünden, Switzerland). Unpublished Ph. D. Thesis, Institute for Mineralogy and Petrography, Basel, Switzerland.
- BRUGGER, J. and BERLEPSCH, P. (1996): Description and crystal structure of fianelite $\text{Mn}_2\text{V}(\text{V}, \text{As})\text{O}_7 \cdot 2\text{H}_2\text{O}$,

- a new mineral from Fianel, Val Ferrera (Graubünden, Switzerland). *Amer. Mineral.*, 81, 1270–1276.
- BRUGGER, J., GIERÉ, R., GRAESER, S. and MEISSER, N. (1997): The crystal chemistry of roméite. *Contr. Mineral. Petrol.*, 127, 136–146.
- CHOPIN, C. (1978): Les paragenèses réduites ou oxydées des concentrations manganésifères des "schistes lustrés" de Haute-Maurienne (Alpes françaises). *Bull. Minéral.* 101, 514–531.
- DE CAPITANI, C. (1983): Petrographische Untersuchungen in der Gegend Furtschellas-Grialetsch (Oberengadin). Diploma thesis Universität Bern.
- DUNN, P.J., PEACOR, D.R., SU, S.-C., NELEN, J.A. and VON KNORRING, J.A. (1986): Johninnesite, a new sodium manganese arsenosilicate from the Kombat mine, Namibia. *Min. Mag.*, 50, 667–670.
- GEIGER, T. and CABALZAR, W. (1989): Ardennit – ein Neufund von der Alp Parsettens, Val d'Err, GR. *Schweizer Strahler*, 8, 5, 201–211.
- GRICE, J.D. and DUNN, P.J. (1994): Johninnesite: Crystal structure determination and its relationship to other arsenosilicates. *Amer. Mineral.* 79, 991–995.
- INNES, J. and CHAPLIN, R.C. (1986): Ore bodies of the Kombat mine, South West Africa-Namibia. In: ANHAESSER, C.R. and MASKE, S. (eds): *Mineral deposits of southern Africa*. Geological Society of Johannesburg, Republic of South Africa, 1789–1805.
- MARCHIG, V., GUNDLACH, H., MOLLER, P. and SCHLEY, F. (1982): Some geochemical indicators for discrimination between diagenetic and hydrothermal metalliferous sediments. *Marine Geology*, 50, 241–256.
- MARQUER, D., CHALLANDES, N. and BAUDIN, T. (in press): Shear zone patterns and strain distribution at the scale of a Penninic nappe: the Suretta nappe (Eastern Swiss Alps). *Journal of Structural Geology*.
- MARTIN-VERNIZZI, S. (1984): La mine de Praborna (Val d'Aoste, Italie): une série manganésifère métamorphisée dans le faciès éclogite. Ph. D. Thesis University Pierre et Marie Curie, Paris, France.
- PEACOR, D.R., DUNN, P.J., SU, S.-C. and INNES, J. (1987): Ribbeite, a polymorph of alleghanyite and a member of the leucophoenicite group from the Kombat mine, Namibia. *Amer. Mineral.*, 72, 213–216.
- PETERS, T., TROMMSDORFF, V. and SOMMERAUER, J. (1980): Progressive metamorphism of manganese carbonate and cherts in the Alps. In: VARENTSOV, I.M. and GRASSELY, G. (eds): *Geology and geochemistry of manganese*. Schweizerbart'sche Verlagsbuchhandlung, Stuttgart, 271–283.
- REINECKE, T. (1986): Phase relationships of sursassite and other Mn-silicates in highly oxidized low-grade, high-pressure metamorphic rocks from Evvia and Andros Island, Greece. *Contr. Mineral. Petrol.*, 94, 110–126.
- ROUSE, R.C., DUNN, P.J. and INNES, J. (1986): Kombatite, the vanadium analogue of sahlinite, from the Kombat mine, South West Africa. *N. Jb. Miner. Mh.*, 519–522.
- SCHMID, S.M., PFIFFNER, O.A. and SCHREURS, G. (1997): Rifting and collision in the Penninic zone of eastern Switzerland. In: PFIFFNER, A. (ed.): *Deep structure of the Swiss Alps – Results from NFP/PNR 20*. Birkhäuser AG, Basel.
- SCHREURS, G. (1991): Structural analysis of the Schams nappes and adjacent tectonic units in the Penninic zone (Grisons, SE-Switzerland). *Mitteilungen aus dem geologischen Institut der Eidgenössischen Technischen Hochschule und der Universität Zürich*, Neue Folge Nr. 283, 204 pp.
- STUCKY, K. (1960): Die Eisen- und Manganerze in der Trias des Val Ferrera. *Beiträge zur Geologie der Schweiz, Geotechnische Serie*, 37. Lieferung, 67 pp.

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